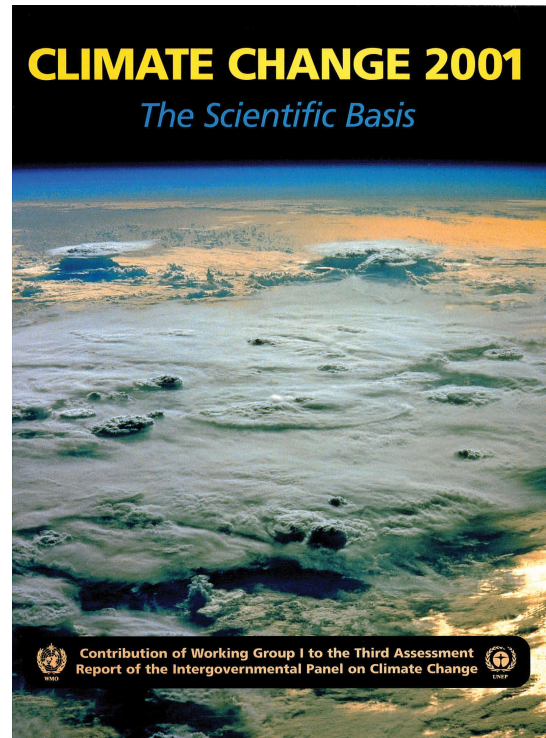


THE CLIMATE CHANGE HOAX UNMASKED



This is the cover of the Intergovernmental Panel on Climate Change (IPCC) Climate Change 2001: The Scientific Basis report.

Table 2.4: Guide to terminology used in paleoclimate studies of the last 150,000 years.

"Event", Stage	Estimated age (calendar years)
Holocene	~10 ky BP to present
Holocene maximum warming (also referred to as "climatic optimum")	Variable?
Last deglaciation	~4.5 to 6 ky BP (Europe) 10 to 6 ky BP (SH)
Termination 1	~18 to 10 ky BP
Younger Dryas	~14 ky BP
Antarctic cold reversal	~12.7 to 11.5 ky BP
Billing-Allerød warm period	14 to 13 ky BP
Last glacial	14.5 to 13 ky BP (Europe)
LGM (last glacial maximum)	~75 to 14 ky BP
Last interglacial peak	~25 to 18 ky BP
Termination 2	~124 ky BP
Bombian/MIS stage 5e	~130 ky BP
Heinrich events	~128 to 118 ky BP
Dansgaard-Oeschger events	Peaks of ice-rafted detritus in marine sediments, ~7 to 10 ky time-scale.
Bond cycles	Warm-cold oscillations determined from ice cores with duration ~2 to 3 ky.
Terminations	A quasi-cycle during the last Ice Age whose period is equal to the time between successive Heinrich events.
	Periods of rapid deglaciation.

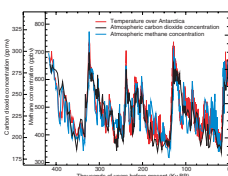


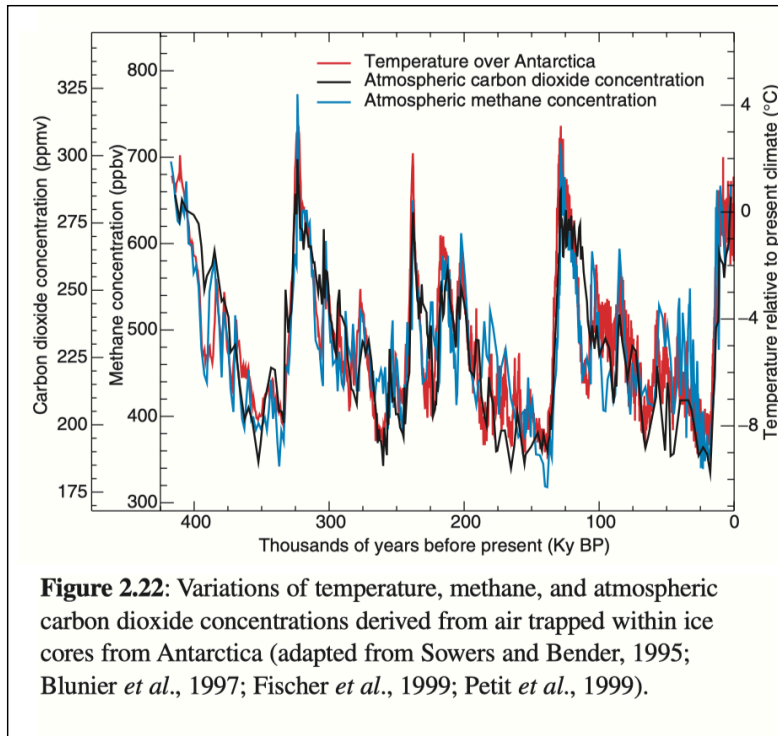
Figure 2.22: Variations of temperature, methane, and atmospheric carbon dioxide concentrations derived from air trapped within ice cores from Antarctica (adapted from Sowers and Bender, 1995; Blunier *et al.*, 1997; Fischer *et al.*, 1999; Petit *et al.*, 1999).

Before reviewing important recent information about rapid changes, we briefly mention progress made on two aspects of the paleoclimate record of relevance for future climate. The first deals with the relationship between modern and past terrestrial data and SSTs around the time of the Last Glacial Maximum (about 20 ky BP); this is important because of the use of glacial data to validate climate models. New results obtained since the SAR both from marine and terrestrial sources (reviewed in Chapter 8), agree on a tropical cooling of about 5°C. The second concerns the greenhouse gas record (CO₂ and CH₄) which has now been considerably extended due to the recent completion of drilling of the Vostok ice

core in central East Antarctica. The strong relationship between CO₂ and CH₄ and Antarctic climate documented over the last climatic cycle has been remarkably confirmed over four climatic cycles, spanning about 420 ky (Figure 2.22). Present day levels of these two important greenhouse gases appear unprecedented during this entire interval (Petit *et al.*, 1999; and Figure 2.22). From a detailed study of the last three glacial terminations in the Vostok ice core, Fischer *et al.* (1999) conclude that CO₂ increases started 600 ± 400 years after the Antarctic warming. However, considering the large uncertainty in the ages of the CO₂ and ice (1,000 years or more if we consider the ice accumulation rate uncertainty), Petit *et al.* (1999) felt it premature to ascertain the sign of the phase relationship between CO₂ and Antarctic temperature at the initiation of the terminations. In any event, CO₂ changes parallel Antarctic temperature changes during deglaciations (Sowers and Bender, 1995; Blunier *et al.*, 1997; Petit *et al.*, 1999). This is consistent with a significant contribution of these greenhouse gases to the glacial-interglacial changes by amplifying the initial orbital forcing (Petit *et al.*, 1999).

We also now have a better knowledge of climate variability over the last few climatic cycles as illustrated by selected paleotemperature records back to about 400 ky (Figure 2.23). The amplitude of the glacial-interglacial temperature change was lower in tropical and equatorial regions (e.g., curve c) than in mid- and high latitudes (other curves). During glacial periods, the climate of the North Atlantic and adjacent regions (curves a and b) was more variable than in the Southern Hemisphere (curve d). Also (not shown), full glacial periods were characterized by very high fluxes of dust (seen in ice-core records and in continental and marine records). A combination of increased dust source area, stronger atmospheric transport and a weaker hydrological cycle (Yang *et al.*, 1996; Makovskii *et al.*, 1999; Petit *et al.*, 1999) probably generated these changes.

This is page 137 which shows Figure 2.22: Variations of temperature, methane, and atmospheric carbon dioxide concentrations derived from air trapped within ice cores from Antarctica (adapted from Sowers and Bender, 1995; Blunier *et al.*, 1997; Fischer *et al.*, 1999; Petit *et al.*, 1999).



This is a close up of the graph on that page. Notice that the graph shows the carbon dioxide levels and the temperature levels rising and falling with each other simultaneously.

That is the current mantra of Climate Change, previously Global Warming, that the increase in carbon dioxide is causing a simultaneous increase in atmospheric temperature.

Notice that the X-axis is in thousands of years. This is a deliberate compression of the original data to make it appear that the increases are simultaneous.

Ice Core Records of Atmospheric CO₂ Around the Last Three Glacial Terminations

Hubertus Fischer, Martin Wahlen, Jesse Smith, Derek Mastroianni, Bruce Deck

Air trapped in bubbles in polar ice cores constitutes an archive for the reconstruction of the global carbon cycle and the relation between greenhouse gases and climate in the past. High-resolution records from Antarctic ice cores show that carbon dioxide concentrations increased by 80 to 100 parts per million by volume 600 ± 400 years after the warming of the last three deglaciations. Despite strongly decreasing temperatures, high carbon dioxide concentrations can be sustained for thousands of years during glaciations; the size of this phase lag is probably connected to the duration of the preceding warm period, which controls the change in land ice coverage and the buildup of the terrestrial biosphere.

Previous studies of Antarctic ice cores (1–3) revealed that atmospheric CO₂ concentrations changed by 80 to 100 parts per million by volume (ppmv) during the last climatic cycle and showed, together with continuous atmospheric measurements (4), that anthropogenic emissions increased CO₂ concentrations from 280 ppmv during preindustrial times to more than 360 ppmv at present, an increase of more than 80% of the glacial-interglacial change. Variations in atmospheric CO₂ concentrations accompanying glacial-interglacial transitions have been attributed to climate-induced changes in the global carbon cycle (5, 6), but they also amplify climate variations by the accompanying greenhouse effect. Accordingly, the relation of temperature and greenhouse gases in the past derived from ice core records has been used to estimate the sensitivity of climate to changes in greenhouse gas concentrations (7) to constrain the prediction of an anthropogenic global warming. This procedure, however, requires the separation of systematic variations representative for all climatic cycles from those specific for each event, as well as a more detailed knowledge of the leads and lags between greenhouse gas concentrations and climate proxies.

To resolve short-term changes in the atmospheric carbon reservoir, to constrain the onset and end of major variations in CO₂ concentrations, and to test whether these variations are temporally representative, we expanded the Antarctic Vostok CO₂ record over the transition from marine isotope stage (MIS) 8 to MIS 7 [about 210 to 250 thousand years (ky) before present (B.P.)] and analyzed the time interval around the penultimate deglaciation (10 to 15 ky B.P.).

Scripps Institution of Oceanography, Geosciences Research Division, University of California San Diego, La Jolla, CA 92093-0220, USA.

mate deglaciation (about 70 to 160 ky B.P.) at a high resolution of 100 to 2000 years (8). This data set was supplemented by a CO₂ record recently derived from the Antarctic Taylor Dome (TD) ice core (6, 9) covering the last 55,000 years. The internal temporal resolution of ice core air samples is restricted by the age distribution of the bubbles caused by the enclosure process (10). This age spread is about 300 years for Vostok (11) and 140 years for the TD ice core (9) at present but about three times higher for glacial conditions (11). The depth-ice age scale used for terminations II and III in the Vostok core is a recently expanded version of the extended glaciological time scale (12). The dating uncertainty (on the order of 10,000 years for termination III) is considerable; however, the absolute time scale is not so important as long as we consistently compare Vostok CO₂ with the Vostok isotope temperature (δD) record.

More important is the relative dating of ice and air at a certain depth. The ice age–air age difference (Δage) was calculated with a climatological firn densification model (13) and varies between about 2000 and 6000 years for warm and cold periods, respectively. The accuracy of the model is better than 100 years for recent periods but on the order of 1000 years for glacial conditions (13), which has to be kept in mind when interpreting the phase shift between ice and gas records of the ice core archive. In the case of termination I, recently published age scales derived by synchronization of CH₄ variations in central Greenland and Antarctic ice cores (13, 14) were used. The precision of the CH₄ correlation is about 200 years for periods of substantial CH₄ change but is not very well constrained in the interval between 17 and 25 ky B.P., when only subtle CH₄ changes occurred. The uncertainty of Δage varies between 100 and 300 years for central Green-

land (13) and between 300 and 600 years for TD (14) during termination I. Further uncertainty is added because the TD CO₂ record has been dated relative to the Greenland Ice Sheet Project 2 (GISP2) core (14), whereas the Byrd and Vostok isotope temperature records have been synchronized with respect to the Greenland Ice Core Project (GRIP) ice core record (15). This uncertainty is not relevant for the interval between 10 and 15 ky B.P., for which dating of GISP2 and GRIP is in good agreement; however, there is a shift of up to 2000 years between the two Greenland reference cores at the age of 20 ky B.P.

In Fig. 1, our data and previously published CO₂ concentration records (1, 6, 8, 11, 12, 16) are compared with Antarctic isotope (temperature) ice core records (13, 17–19). Note that the CO₂ concentrations represent essentially a global signal. In contrast, the geographical representativeness of isotope temperature records may vary from a synoptical to hemispherical scale and accordingly within different cores with increasing variability for shorter time scales. The Vostok and TD CO₂ data presented here are in good agreement with previous CO₂ values. On a 10,000-year time scale, CO₂ covaries with the isotope temperatures with minimum glacial CO₂ concentrations of 180 to 200 ppmv glacial-interglacial transitions accompanied by a rapid increase in CO₂ concentrations to a maximum of 270 to 300 ppmv, and a gradual return to low CO₂ values during glaciation. On a shorter time scale, however, a much more complex picture evolves.

The onset of the atmospheric CO₂ increase during termination I recorded in the TD record is at 19 to 20 ky B.P. The rise in the long-term trend in CO₂ concentrations seems to be about 1000 years earlier than the rise in Vostok δD values. In contrast, temperatures apparently started to rise at 20 ky B.P., as recorded in the Antarctic Byrd and the Greenland GRIP ice core (13). Again, CO₂ concentrations in the Byrd record increase ~2000 ± 500 years later than those in the TD data. In view of the excellent agreement for the rest of the CO₂ records, these discrepancies can be attributed to the insufficient age constraint during the onset of termination I induced by the different Greenland reference cores. No such dating uncertainties are encountered for the interval between 10 and 15 ky B.P. Maximum CO₂ concentrations of 270 ppmv are reached at 10.5 ky B.P. (9, 600 to 1000 years after the isotope temperature maximum in the Byrd record (20)). The CO₂ peak is followed by a decrease of 5 to 10 ppmv until 8 ky B.P., after which CO₂ concentrations gradually rise to the preindustrial value of 280 ppmv (9). A delay in the increase of CO₂ concentrations with respect to the warming during deglaciation is also indicated by a brief 10-ppmv decline in CO₂ concentrations

Downloaded from www.sciencemag.org on March 20, 2012

This is the Fischer, *et al*, 1999 paper that was referenced in figure 2.22.

REPORTS

found in seven samples during the interval 14 to 11 ky B.P. This distinct feature lags the Antarctic Cold Reversal (ACR) in the Antarctic isotope temperatures (δT) by 300 to 500 years but occurs 1000 years before the Younger Dryas cooling event.

A dip in CO_2 concentrations at 135 ky B.P. precedes the start of the increase in CO_2 concentrations during termination II, which reaches a maximum of 280 ppmv at 125 ky B.P. (Like in the Holocene, CO_2 concentrations decrease after the initial maximum to ~275 ppmv. The onset of the major warming during termination II is hard to define, but during the penultimate warm period, CO_2 concentrations reach their maximum 400 ± 200 years later than Antarctic temperatures. In the following 15,000 years of the Eemian warm period, CO_2 concentrations do not show a substantial change despite distinct cooling over the Antarctic sea sheet. Not until 6000 years after the major cooling in MIS 5.4 does a substantial decline in CO_2 concentration occur. Another 4000 to 6000 years is required to return to an approximate in-phase relation of CO_2 with the temperature variations.

Finally, termination III starts with a CO_2 concentration of 265 ppmv at 244 ky B.P., slightly higher than that for the beginnings of terminations I and II. At that time, temperatures had already increased since the glacial temperature minimum at ~260 ky B.P. CO_2 concentrations rise slowly from 244 to 241 ky B.P. and then rapidly to more than 300 ppmv at 238 ky B.P. Keeping the rather coarse resolution of the δD record before 238 ky B.P. in mind, the major increase in CO_2 tends to lag temperature during the transition, reaching a maximum CO_2 concentration 600 \pm 200 years after the peak in δD . In contrast to the case for the Eemian, high CO_2 concentrations are not sustained during MIS 7 but follow the rapid temperature drop into MIS 7.4. Minimum CO_2 concentrations as low as 210 ppmv are reached 1000 to 2000 years after the minima in isotope temperature during MIS 7.4. A short, warm event during the mid-glacial interval at 224 to 228 ky B.P. appears to be reflected in a 30-ppmv increase in atmospheric CO_2 concentrations with a phase lag of about 1000 \pm 600 years relative to temperature. Another warm event at the beginning of the warm period MIS 7.3 is accompanied by a 10-ppmv increase in CO_2 concentration, which appears to be in phase with the temperature record. The variations in CO_2 concentrations during these events are much larger than anticipated from the Vostok isotope temperature changes and do not have any counterparts during MIS 5.

Comparison of the sequence of events for the three time intervals described above suggests that the carbon cycle-climate relation should be separated into (at least) a deglaciation and a glacialation mode. Atmospheric CO_2 concentrations show a similar increase for all

three terminations, connected to a climate-driven net transfer of carbon from the ocean to the atmosphere (6). The time lag of the rise in CO_2 concentrations with respect to temperature change is on the order of 400 to 1000 years during all three glacial-interglacial transitions. Considering the uncertainties in lags (between 100 and 1000 years for recent and glacial conditions), such a lag can still be explained by an overestimation of lags for glacial conditions. The good agreement of the lags model with the measured values for the present supports the idea that at least the lag at the beginning of the warm periods is real. The size of this lag is on the order of the ocean mixing time (for a well-ventilated ocean like today), which is the major control for the time constant of equilibration within the deep ocean-atmosphere carbon system after climate-induced changes. In the case of a recent anthropogenic warming, the exter-

nal climate forcing by CO_2 emissions due to combustion of fossil fuel leads climate variations, so the application of the CO_2 -climate relation deduced from the past on a recent global warming seems not to be straightforward.

The situation is even more complicated for the interglacial and glacial periods. During the extended Holocene and Eemian warm periods, atmospheric CO_2 concentrations drop by ~10 ppmv after an initial maximum, attributable to a substantial increase in the terrestrial biospheric carbon storage extracting CO_2 from the atmosphere. In the case of the Eemian, CO_2 concentrations remain constant after the initial maximum in MIS 5.5 despite slowly decreasing temperatures; during the Holocene, atmospheric CO_2 concentrations even increase during the last 8000 years. Application of a carbon cycle model to CO_2 and $\delta^{18}\text{O}$ ice core data for the Holo-

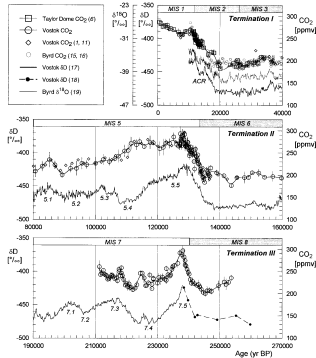


Fig. 1. Records of atmospheric CO_2 concentrations and isotope temperature records derived from the Antarctic Byrd, Vostok, and TD ice cores during the deglaciation and glacialation events around the last three glacial terminations. Error bars in CO_2 concentration data represent 1 σ of replicate measurements at the same depth interval. The long-term trend in CO_2 concentrations is indicated by a cubic spline approximation ($F = 5 \times 10^{-12}$ of our data set). For convenience, marine isotope stages (22) are indicated as referred to in the text.

This is the second page of the paper. The high lighted sentence in the second column reads as follows.

“The time lag of the rise in CO_2 (carbon dioxide) concentrations with respect to temperature change is on the order of 400 to 1000 years during all three glacial-interglacial transitions.”

The temperature change comes before the rise in carbon dioxide by hundreds of years, not the other way around.

The deception is the compression of the graph in the report to make it look like it is simultaneous when it is not.